

Evaluation of a Cloud Resolving Model Using TRMM Observations for Multiscale Modeling Applications

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Popular Summary

As an intermediate step between conventional general circulation models (GCMs) and global cloud resolving models (CRMs), many climate prediction centers are embedding a CRM in each grid cell of a conventional GCM. These Multiscale Modeling Frameworks (MMFs) represent a theoretical advance over the use of conventional GCM cloud and convection parameterizations in that they directly simulate a distribution of convective elements in a given GCM grid box. Though they represent a theoretical advance over conventional GCMs, MMFs have been shown to exhibit an overproduction of precipitation in the tropics during the northern hemisphere summer. This precipitation bias is particularly pronounced over the Tropical Western Pacific and South China Sea during the Asian Monsoon.

In this study, the CRM component of the NASA Goddard MMF is evaluated against retrievals derived from multiple instruments aboard the Tropical Rainfall Measuring Mission (TRMM) satellite platform. To examine the CRM in isolation from the parent GCM, the NASA Goddard Cumulus Ensemble (GCE) model is driven with observed large-scale forcing derived from soundings taken during the South China Sea Monsoon Experiment. Simulations of clouds, precipitation, and radiation over the South China Sea are then compared with TRMM retrieved clouds, precipitation, and radiative fluxes. It is found that the GCE configuration used in the NASA Goddard MMF responds too vigorously to the imposed large-scale forcing, accumulating too much moisture and producing too much cloud cover during convective phases, and overdrying the atmosphere and suppressing clouds during monsoon break periods.

Sensitivity experiments reveal that changes to microphysical parameters that determine the precipitating (snow and graupel) ice particle size distribution have a relatively large effect on simulated clouds, precipitation, and radiation, while changes to grid spacing and domain length have little effect on simulation results. The results motivate a more detailed and quantitative exploration of the sources and magnitude of the uncertainty associated with specified cloud microphysical parameters in the CRM components of MMFs. Data assimilation experiments, in which this uncertainty is quantified for each microphysical parameter, as well as comparisons between microphysical schemes with differing levels of complexity, are underway.

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Abstract

The climate change simulation community is moving toward use of global cloud resolving models (CRMs), however, current computational resources are not sufficient to run global CRMs over the hundreds of years necessary to produce climate change estimates. As an intermediate step between conventional general circulation models (GCMs) and global CRMs, many climate analysis centers are embedding a CRM in each grid cell of a conventional GCM. These Multiscale Modeling Frameworks (MMFs) represent a theoretical advance over the use of conventional GCM cloud and convection parameterizations, but have been shown to exhibit an overproduction of precipitation in the tropics during the northern hemisphere summer.

In this study, simulations of clouds, precipitation, and radiation over the South China Sea using the CRM component of the NASA Goddard MMF are evaluated using retrievals derived from the instruments aboard the Tropical Rainfall Measuring Mission (TRMM) satellite platform for a 46-day time period that spans 5 May - 20 June 1998. The NASA Goddard Cumulus Ensemble (GCE) model is forced with observed large-scale forcing derived from soundings taken during the intensive observing period of the South China Sea Monsoon Experiment. It is found that the GCE configuration used in the NASA Goddard MMF responds too vigorously to the imposed large-scale forcing, accumulating too much moisture and producing too much cloud cover during convective phases, and overdrying the atmosphere and suppressing clouds during monsoon break periods. Sensitivity experiments reveal that changes to ice cloud microphysical parameters have a relatively large effect on simulated clouds, precipitation, and radiation, while changes to grid spacing and domain length have little effect on simulation results. The results motivate a more detailed and quantitative exploration of the sources and magnitude of the uncertainty associated with specified cloud microphysical parameters in the CRM components of MMFs.

1 Introduction

One of the greatest sources of uncertainty in current estimates of climate system response to surface warming is the influence of clouds and precipitation. Specifically, it is unknown whether, given a globally averaged rise in surface temperature, clouds will act to enhance or mitigate the climate system response. Of particular interest is deep convection in the tropics, which plays an important role as it links the fluxes of short and longwave radiation with the large scale circulation and provides an important mid-tropospheric energy source through the release of latent heat. Though the effects of deep convection and the associated cloud field are felt on scales greater than 1000 km, the processes that critically determine the amount and intensity of precipitation, as well as the properties and extent of upper-level cirrus operate on the scales of a few km. It is in part due to this intrinsic separation in scales that it has been traditionally difficult for General Circulation Models (GCMs) to accurately simulate the observed distribution of clouds and precipitation in the tropics (Bony et al. 2006, Soden and Held 2006).

Clouds and precipitation have typically been parameterized in climate models by assuming a distribution of cloud heights and convective cores within a single GCM grid cell (Arakawa and Schubert 1974, Zhang and McFarlane 1995, Sud and Walker 1999). This type of representation suffers from uncertainty in the inherent assumptions about the characteristics of clouds and precipitation on the sub-GCM grid scale, leading to a wide range of uncertainty in the interaction between clouds and radiation, as well as in the vertical distribution of latent heat release. After three decades of concentrated efforts designed to develop, evaluate, and improve conventional GCM cloud and convective parameterizations, it is generally acknowledged that clouds, precipitation, and their interaction with radiation can be most realistically simulated using models that are run on the scales of the cloud processes themselves (grid lengths of at most four kilometers). Though computational resources are continually increasing, it is still not possible to perform global simulations of climate change on cloud-resolving scales. As an intermediate step between GCMs and global cloud resolv-

ing models, selected centers have implemented so-called “Multi-scale Modeling Frameworks” (MMFs). These models employ a Cloud Resolving Convective Parameterization (CRCP, Grabowski 2001) or “Super-Parameterization” (Khairoutdinov and Randall 2001, Randall et al. 2003), and use a cloud resolving model to represent cloud-scale processes by embedding a small two-dimensional CRM in each GCM grid cell. Though certain features of the observed climate system (e.g. the Madden Julian Oscillation) are more realistically simulated in MMFs, there are problems in MMFs that are not observed in conventional GCMs. The most prominent example is the existence of the so-called “Great Red Spot”; a region of anomalously large precipitation centered over the South China Sea and Bay of Bengal during the northern hemisphere summer (Khairoutdinov et al. 2005, Tao et al. 2007). Sensitivity tests have indicated that the use of a small three-dimensional domain, and the adjustment of 2D CRM orientation within each GCM grid cell may help to reduce this anomaly, and Luo and Stephens (2006) have hypothesized that the problem arises from a convection-wind-evaporation feedback operating on a small cyclic CRM domain. In general, the problem illustrates the need for a more systematic evaluation of the CRM component of a multiscale modeling framework. In particular, it remains to be demonstrated that the CRM simulated distribution of clouds and precipitation and the associated interaction with visible and infrared radiative fluxes is consistent with observations.

This paper addresses two fundamental questions regarding use of a CRM as a replacement for the conventional convective parameterization in a GCM. First, given the limited domain and two-dimensional nature of CRM simulations in a MMF, can the CRM correctly reproduce the observed statistics of clouds, precipitation, and radiation? Second, what are the dominant sources of uncertainty in the CRM, and can observations be brought to bear to reduce this uncertainty? The first question is addressed by comparing statistics of the clouds, precipitation, and radiation produced by the NASA Goddard Cumulus Ensemble (GCE) cloud-resolving model used in the NASA Goddard Space Flight Center MMF to retrievals of clouds, precipitation, and radiation from the Tropical Rainfall Measuring Mission

(TRMM) satellite platform. A partial answer to the second question is provided through a set of sensitivity experiments, in which changes to domain size, grid spacing, and cloud microphysical assumptions are applied to the model and the resulting fields are compared to a control simulation.

A multiscale modeling framework typically consists of a two-way interaction between CRM and GCM, in which GCM “large-scale” tendencies of temperature and water vapor are first applied over the a CRM simulation time equal to a GCM timestep, after which the CRM tendencies of temperature and water vapor are applied on the GCM grid. The resulting CRM fields are therefore affected by errors in both CRM and GCM. To evaluate clouds, precipitation, and radiation produced by the CRM apart from uncertainties in the parent GCM, CRM simulations are performed in the presence of observed large-scale forcing computed from a sounding network deployed during the South China Sea Monsoon Experiment (SCSMEX) that ran from 5 May through 20 June, 1998. Large-scale forcing includes the effects of advection of temperature and water vapor, as well as large-scale tendencies of the three-dimensional wind. A thorough description of the methodology used to generate the forcing dataset is contained in (Johnson and Ciesielski 2000). Forcing the GCM with observed large scale fields is analogous to embedding the CRM within a GCM that perfectly simulates the large-scale flow, with the primary difference being that this is a purely one-way interaction with no feedback allowed from the CRM back to the large scales. When forced with observed large-scale advective tendencies of wind, temperature, and moisture, the CRM should produce clouds and precipitation that are consistent with observed conditions, and ideally, consistent with the observed distribution of clouds and precipitation.

The comparison of model with observations is done in a statistical manner since it is the temporal and spatial distribution of clouds and cloud properties that have an effect on key climate variables. As such, it is not necessary for the model to reproduce each observed convective element or system at the exact time and place it occurred in reality. Instead, it is sufficient that the model produce the appropriate distribution of clouds and precipitation and

that the effect of clouds on radiative fluxes and heating rates is consistent with observations. To this end, statistics of GCE-simulated clouds, precipitation, and radiation, accumulated over the SCSMEX field experiment are compared to TRMM observations via two quantitative metrics. The first is a measure of the center of mass of the distribution, which can be defined as the mean, median, or mode, depending on the specifics of the comparison. The second is an integrated measure of the difference between each histogram over the combined range of values found in observations and model. This measure is effectively the sum of differences between histograms in each bin, and can be computed as the absolute value (integrated absolute difference, IAD) or alternatively as the root mean square (integrated RMS). The IAD carries the added benefit of an intuitive interpretation, as a 50% difference in PDF mass is computed as an IAD of 1.0, while a perfect mismatch (no shared mass between histograms) is computed as an IAD of 2.0. An illustration of the utility of both measures can be found in figure 1, in which a comparison of two idealized histogram is presented. While the mean of the two PDFs is identical in the first case, the IAD reveals a difference in structure. The second case has the identical IAD as the first case, but the means are shifted, indicating that the difference in PDF may be due more to the fact that there is a bias in the solution than to a structural difference in PDF.

Two time periods are examined in this study: a 46-day interval, equal to one full TRMM precession cycle, and a shorter 9-day period during which the SCSMEX domain was characterized by repeated development of convective squall lines evolving in the presence of strong vertical wind shear. This second case is of particular interest for two reasons (1) it allows the evaluation of CRM simulations of convection in the region of the Great Red Spot and (2) in past numerical studies of convection it has been demonstrated that strongly sheared convective squall lines are more likely to be realistically represented in the two-dimensional framework than is convection that develops in weak shear or under suppressed conditions (Grabowski et al. 1998, Tompkins 2000, Petch and Gray 2001). Comparisons between model data and observations will be used to demonstrate that the GCE develops a moist bias over

the course of the 46-day integration, and that this bias exists independent of changes to grid length or spacing. The steady accumulation of tropospheric water vapor leads to overprediction of cloud fraction, more frequent and intense precipitation, liquid and ice cloud that is thicker than observed, and cloud radiative forcing that is too strong. It will also be shown that modification of cloud microphysical parameters can lead to significant changes to the statistics of all simulated fields, and potentially to an improved agreement with observations. The results imply that the details of the cloud microphysical parameterization play a key role in determining the statistics of clouds and precipitation simulated by the CRM, and that changes in grid spacing or geometry cannot eliminate problems associated with uncertainties in the cloud microphysical scheme. It is clear from this study that prior to use of a CRM in an MMF, uncertainties in the cloud microphysical representation must be quantified and mitigated.

The remainder of this paper is organized as follows. A brief overview of the SCSMEX field campaign and computed large-scale forcing fields, along with a description of the TRMM multisensor retrieval algorithm, retrieved products, and estimated errors is presented in section 2. The details of the GCE model and of the specific configuration used in the GSFC MMF are described in section 3. Results of the GCE comparison with TRMM are presented in section 4, while sensitivity to grid and cloud microphysical parameters is explored in section 5. Summary, conclusions, and suggestions for future work are offered in section 6.

2 TRMM Retrievals and Description of SCSMEX

2.1 South China Sea Monsoon Experiment

The South China Sea Monsoon Experiment (SCSMEX, Lau et al. 2000) was conducted during May and June 1998 to examine the mechanisms associated with the onset of monsoon convection over the northern and southern South China Sea (SCS). It also served as a validation campaign for the newly-launched TRMM satellite (Kummerow et al. 2000).

During SCSMEX, radiosondes were regularly launched from several sites on and around the SCS (Ding and Lau 2001) with frequencies of four times daily over each of two Enhanced Sounding Arrays; one centered around the TOGA-BMRC dual-doppler radar network in the northern SCS, and the other centered around the *Kexue #1* GPS sounding site. Gridded fields of temperature, specific humidity, geopotential height, and the horizontal components of the wind were produced over the South China Sea at one-degree resolution using a multi-quadric interpolation scheme (Nuss and Titley 1994, Johnson and Ciesielski 2000). These fields were subsequently used to compute the pressure vertical velocity (ω) and large-scale advective forcing of temperature and water vapor at six-hourly intervals averaged over both the Northern Enhanced Sounding Array (NESA) and Southern Enhanced Sounding Array (SESA) regions. These forcing fields have been used to drive 2D GCE simulations in previous studies of the Asian monsoon (Tao et al. 2003a, Tao et al. 2004). Depiction of the SCS with NESA and SESA regions is provided in figure 2; only results from the NESA will be used in this paper.

Timeseries of TRMM observed daily domain-averaged precipitation rate and outgoing longwave radiation (OLR) over the NESA region (Fig. 3) reveals convection occurring in two distinct phases; a pre- and during-monsoon onset period that lasted from 18-26 May, and a post-onset episode that lasted from 2-11 June. Examination of the mean vertical shear for each of these periods (not shown) reveals weak unidirectional shear (approximately 10 m s^{-1}) that extends through the depth of the troposphere for the May time period, with stronger shear (approximately 15 m s^{-1}) during 2-11 June that reverses sign in the mid-troposphere. Consistent with the expectation of more strongly organized propagating convection under conditions of strong low-level vertical shear (Fovell and Ogura 1989), convection was observed to be more vigorous during the June time period with nearly double the mean precipitation rate observed during the pre-onset period (Johnson and Ciesielski 2002). OLR is generally anti-correlated with the precipitation rate, with relatively large values of OLR observed during periods in which precipitation was light, and small OLR observed during periods of

heavy precipitation. Consistent with the production of stratiform precipitation regions and the persistence of clouds following a period of deep convection, the timing of maxima in OLR lag the peak precipitation slightly.

2.2 TRMM Multisensor Retrieval Algorithm

Observations used in this study derive from a multisensor retrieval algorithm that combines ice cloud information from the Visible and Infrared Scanner (VIRS) (Cooper et al. 2003), liquid cloud information from the TRMM Microwave Imager (TMI) (Greenwald et al. 1993), and precipitation information from the TMI-based Goddard Profiling algorithm (GPROF) (Kummerow et al. 2000) to produce an analysis of the precipitable water vapor (PWV), precipitation rate, liquid water path (LWP), ice water path (IWP), and liquid and ice cloud fraction. Visible and infrared top of atmosphere fluxes and shortwave and longwave heating and cooling rates computed in three discrete layers; low (0.5 - 2.5 km), middle (2.5 - 5.0 km), and high (5.0 - 17.0 km) (L'Ecuyer and Stephens 2003) are obtained from broadband radiative transfer calculations in each pixel. The estimated uncertainties in all retrieved parameters (Table 1) are based on a combination of rigorous sensitivity studies and product intercomparisons (L'Ecuyer and Stephens 2002, Cooper et al. 2003, L'Ecuyer and Stephens 2003). The time period that spans 0000 UTC 5 May - 0000 UTC 20 June 1998 was specifically chosen to match a single 46-day TRMM precession period,¹ the time interval over which TRMM repeats an observation of the identical position on the Earth's surface at the identical local time. Use of the full precession period helps avoid a bias toward observing at a particular local time, and consequently avoids biases in observations of the diurnal cycle of clouds and precipitation.

¹See <http://tsdis.gsfc.nasa.gov/overflight/PredictLocalSolar.html> for more details.

3 Description of the NASA Goddard Cumulus Ensemble Model

The NASA Goddard Cumulus Ensemble model is a nonhydrostatic cloud system resolving model based on the work of Soong and Ogura (1980), Soong and Tao (1980), and Tao and Soong (1986). It employs either an anelastic or compressible solution to the atmospheric governing equations, and can be run in either two (one vertical and one horizontal) or three spatial dimensions. The model prognostic variables include potential temperature, perturbation pressure, turbulent kinetic energy, and the three Cartesian velocity components, as well as water vapor mass mixing ratio and the mixing ratios of two liquid and three ice condensate species. A fourth order accurate scheme is used to advect velocity components, while a multidimensional positive definite advection transport algorithm (Smolarkiewicz 1983, 1984, Smolarkiewicz and Grabowski 1990) is used to advect all scalar quantities. Forward time differencing is employed for scalar variables, while a time splitting leapfrog scheme is used for the velocity.

The subgrid turbulence scheme is based on Deardorff (1975), Klemp and Wilhelmson (1978), and Soong and Ogura (1980), and includes the effect of condensation on the generation of subgrid-scale kinetic energy. Interaction of visible and infrared radiation with atmospheric trace gases, clouds, and the surface is parameterized according to Chou and Kouvaris (1991), Chou et al. (1995), Kratz et al. (1998), Chou et al. (1999), and Chou and Suarez (1999). Infrared fluxes are computed via a k-distribution method, while the solar radiation scheme uses a delta-Eddington approximation to compute two-stream radiative fluxes, and accounts for absorption by water vapor, CO_2 , O_2 , and O_3 , as well as the details of the cloud and ice particle size distribution. Surface fluxes are computed using the TOGA-COARE bulk flux algorithm (Fairall et al. 1996), and a 5 km deep Rayleigh relaxation (absorbing) layer is used at the top of the model to prevent reflection of vertically propagating gravity waves.

The GCE configuration employed in the NASA Goddard multi-scale modeling framework (Tao et al. 2007) employs a 256 km horizontal domain with 4 km grid spacing and cyclic boundary conditions in the horizontal. 29 vertical levels are used, which stretch in spacing from 80 meters near the surface to approximately 1000 meters in the upper troposphere and lower stratosphere, and the model top is located at 23 km. In this study, the model is forced at every grid point and at every timestep with the SCSMEX large-scale advective and thermodynamic tendencies of temperature and water vapor, as well as with large-scale horizontal winds. The GCE cloud microphysical scheme is derived from a combination of Lin et al. 1983 and Rutledge and Hobbs (1984), with updates that include a sequential saturation adjustment scheme (Tao et al. 2003a), and adjustments to the Bergeron process, dry growth of graupel, and ice sedimentation according to Lang et al. (2007). This is a single-moment bulk scheme, which predicts the mass mixing ratio of each of two liquid and three ice species, and assumes monodisperse cloud droplets and pristine ice crystals. Precipitating species (rain, snow, and graupel) are assumed to follow a Marshall-Palmer particle size distribution with fixed slope intercept and particle density.

4 Results

Comparison between GCE simulations and TRMM retrievals is performed over two different time periods: simulated fields are first compared with TRMM observations over the full 46-day observation period, then the focus of the comparison is narrowed to the nine-day period spanning 2-11 June. The full 46-day integration includes both periods of active convection and periods during which convection is suppressed, allowing for an assessment of the model performance under a range of different conditions. By contrast, the 2-11 June case consistently exhibits convection over the entire time period (Fig. 3), which accommodates inspection of model performance during conditions for which the MMF was originally designed. In each case, GCE output aggregated to the size of TRMM retrieval pixels are

matched to a time window within ± 30 minutes of each TRMM overpass. Examination of the statistics accumulated over the 46-day integration when all GCE timesteps are used (not shown) indicated differences of at most 5% from the statistics when only TRMM-matched times were used. In contrast to the 46-day integration, 2-11 June TRMM observations of the SCSMEX domain are not evenly distributed in time. This leads to a bias in the diurnal cycle when all GCE output times are used in the comparison and to a corresponding difference in statistics. For consistency, and in order to avoid any bias during the 2-11 June time period, only the TRMM-matched times are used in the comparisons documented below.

4.1 5 May - 20 June 1998

Histograms of simulated and observed fields from the full 46-day integration are plotted in figure 4, while comparison of histogram means and integrated absolute difference (IAD) are presented in tables 4 and 5, respectively. The PDF of simulated precipitable water vapor (Fig. 4a) exhibits a bimodal distribution, with a dry peak in the observations centered between 50 and 52.5 kg m^{-2} and a moist peak centered between 60 and 62.5 kg m^{-2} . The GCE also produces a bimodal distribution, however, the two peaks are well separated from each other; the dry peak is far too dry at 32.5 kg m^{-2} , and the moist peak is too moist at 70-72.5 kg m^{-2} . The implication is that the GCE response to imposed forcing may be too strong, with too much drying during suppressed conditions, and too much moistening during periods of active convection. Liquid and ice water path statistics appear to compare quite favorably over the full integration (Figs. 4b and 4c), with the sole discrepancy being an overprediction of thick clouds; clouds with LWP and IWP greater than 5 kg m^{-2} . Comparison of cloud fraction observed by TRMM and simulated by the GCE model (Table 2) reveals that 51% of total GCE pixels were cloud-covered as compared with only 42% of the TRMM pixels. Examination of the PDF of precipitation rate (Fig. 4d) indicates that the GCE produces too much intense precipitation (rain rate greater than 10 mm h^{-1} compared with the observations, and too little light precipitation (rain rate less than 10 mm h^{-1}). This is

reflected in the fact that only 5% of the GCE pixels had a precipitation rate greater than the assumed 1 mm h^{-1} detection threshold as compared with 13% of TRMM pixels (Table 2).

PWV, LWP, IWP, and precipitation serve as effective illustrations of the utility of both the mean and integrated absolute difference in comparing observed and simulated distributions. In particular, comparison of the means of retrieved and simulated PWV indicate a relatively close fit between the two distributions, as there is approximately a 10% difference in means. By contrast, the integrated absolute difference between the two PDFs is 1.28; a greater than 100% disagreement between the PDFs, indicating that there are significant differences in the structure of the two histograms. The LWP, IWP, and precipitation rate histograms exhibit very different means, but the IAD is quite low (less than 0.1 for LWP and IWP, and less than 0.4 for precipitation rate); an indication that though the overall structure of the PDFs agrees quite well, there is an error due to skewness. In general, the combination of comparable means with large (>1.0) IAD indicates either a misrepresentation of the dispersion or multiple modes in the observed or simulated PDF. By contrast, if the IAD is low (< 0.5), but the mean is displaced (as in LWP, IWP, and precipitation rate), it is an indication that the general structure of the PDF is well-represented, but that the skewness is not properly represented in the simulated PDF.

The effects of clouds on radiation can be assessed through subtraction of the clear-sky from all-sky radiative fluxes and heating rates; the resulting quantities are hereafter referred to as *cloud radiative forcing* and *cloud-affected* heating rates, respectively.² Examination of the cloud radiative forcing (Figs. 4e and 4f) reveals an overprediction of longwave and shortwave cloud forcing in the GCE simulations; the GCE LW cloud forcing histogram has two peaks, as in the observations, however, both the low and high peaks are biased toward

²Subtraction of clear-sky fluxes and heating rates serves to filter any bias in the comparison that might be due to use of temperature, water vapor, and gaseous absorption profiles in the TRMM retrievals that differ from those simulated by the model. Since the veracity of simulated water vapor can be assessed through comparison of GCE and TRMM PWV, and since it is the cloud effect on the radiation field that is the focus of the use of a CRM in a MMF, comparison of only the cloud-affected component of fluxes and heating rates is appropriate in this context.

large absolute values of radiative forcing. The peak in the simulated shortwave cloud forcing is located in the range between -0 to -25 W m^{-2} , which is close to the observed -25 to -50 W m^{-2} , but the model produces too many instances of large negative shortwave cloud radiative forcing relative to the observations. Both long and shortwave results are consistent with overprediction of thick clouds in the GCE, and cloud-affected heating rates follow suit. In particular, at low levels (Fig. 4g), the GCE PDF is missing the enhanced longwave cloud-induced cooling, and produces cloudy heating that is too weak by approximately $1.0\text{-}1.5 \text{ K d}^{-1}$. At mid-levels (Fig. 4h), the GCE approximately reproduces the observed histogram of longwave cooling, but overpredicts cloudy heating, while at high levels (Fig. 4i), the GCE exhibits a bias toward greater cooling than in the observations. The lack of low-level longwave cooling is consistent with underprediction of low cloud under otherwise clear skies; in the case of low cloud under high cloud, cloudy cooling is reduced. Underprediction of mid-level cloudy cooling is consistent with the presence of too much high cloud in the GCE, as most of the midlevel cloud tends to also be associated with the presence of high cloud (not shown). Simulated shortwave cloud-affected heating rates are biased toward too much cooling at all levels, indicating an overprediction of shortwave reflection and consequently clouds that are too optically thick in the visible. It is interesting to note that, although the form of the PDFs for low and mid-level cloudy heating are quite similar (Figs. 4k and 4l), the PDF of simulated shortwave cloudy heating at upper levels appears to be missing a distinct mode at small cooling rates. This, combined with results from low and mid levels provides a further indication that simulated clouds are too thick relative to those that are observed.

4.2 2-11 June 1998

A comparison of histograms of observed fields from TRMM with GCE output over the 2-11 June time period is presented in figure 5. Though a double peak in the PWV field is still clearly visible in the TRMM observations, the moist mode located between 60-65 kg

m^{-2} is now the primary one. By contrast, the GCE PWV has only a single primary peak, located at $70\text{-}75 \text{ kg m}^{-2}$. Comparison of precipitable water vapor from the June time period with PWV from the entire integration indicates that the peak PWV for the June case is collocated with the GCE moist mode in the full integration, and that it is still biased moist compared with the TRMM retrievals. This is reflected in both the mean and IAD; mean simulated PWV is now significantly different from the retrievals, while the IAD indicates a disagreement of nearly 100% ($\text{IAD} = 2.0$). Compared with the full integration, the June case exhibits relatively greater departures from TRMM in the liquid and ice water paths; both simulated LWP and IWP are missing the secondary peak at LWP of $1.5\text{-}2.0 \text{ kg m}^{-2}$ and IWP of $2.0\text{-}2.5 \text{ kg m}^{-2}$, and both exhibit more occurrences of LWP and IWP greater than 5 kg m^{-2} compared with the observations. The histogram mean and IAD again reflect an increased departure from observations in the 2-11 June case; IAD, in particular, is twice that computed for histograms over the full 46-day period. In comparing simulated and observed LWP and IWP, it should be noted that the differences in the means are quite small relative to the range of simulated values, and that the IAD values reflect a less than 10% departure from observations for the June case, and less than 5% departure for the full 46-day period. By contrast, the difference in mean precipitation rate reflects a significant shift toward higher rain rates in the model, and IAD values indicate a greater than 25% difference in the PDFs for 2-11 June; an indication that, though the integrated condensate mass may be well-represented in the GCE, the convection is too intense.

LW cloud radiative forcing in the GCE exhibits a nearly flat histogram from $30\text{-}200 \text{ W m}^{-2}$, while the TRMM retrievals peak noticeably at around 80 W m^{-2} , with a steady decrease to 200 W m^{-2} . Comparison of simulated and retrieved shortwave cloud radiative forcing again indicates a preponderance of highly reflective cloud in the GCE model, as the peak at low values is much reduced in the GCE, and the simulated PDF exhibits a pronounced mode at cloud radiative forcing values less than -400 W m^{-2} . By contrast, the observations exhibit a sharp decrease in the occurrence of shortwave cloud radiative forcing

less than -100 W m^{-2} and almost zero occurrence of values less than -400 W m^{-2} . Longwave cloud affected heating rates exhibit the structures that are similar to those in the long-term integration, though there is a relatively greater bias toward higher longwave heating at low and middle levels, and toward smaller heating and larger cooling at upper levels. Shortwave cloud-affected heating in the model is, as in the long-term integration, biased toward stronger rates of cooling, though it is interesting to note that the PDFs in the June case are uniformly narrower than in the 46-day period. In addition, the secondary and tertiary modes in the observed upper-level shortwave cloudy heating are greatly diminished relative to the full time period. This is most likely a reflection of the fact that the active phase of the monsoon is characterized by cirrus anvils in the domain with far fewer middle and low clouds and less thin cirrus.

The general picture that emerges from comparison of the control GCE simulation with TRMM retrievals is a model that responds too vigorously to the imposed large-scale forcing, accumulating too much moisture and producing too much cloud cover during convective phases, and overdrying the atmosphere and suppressing clouds during monsoon break periods. Evidence for this is provided in figure 6, in which the daily mean precipitable water vapor timeseries is plotted from the GCE control run, the TRMM retrievals, and the SC-SMEX forcing dataset. From this plot, it can be noted that, though the PWV in the forcing dataset is consistently higher than that in the TRMM retrievals by a few kg m^{-2} , it tracks the observed temporal changes in integrated water vapor with great fidelity. In contrast, the GCE simulated water vapor decreases dramatically early in the time period, and increases unrealistically in the latter half. It is interesting to note that, despite the model's problems simulating the observed distribution of water vapor, the distribution of integrated liquid and ice cloud mass is generally well-simulated by the GCE, indicating that the cloud microphysical scheme is producing physically realistic clouds. The fact that the precipitation rate is generally overpredicted compared with observations indicates overactive convection in the model, while the discrepancies in the cloudy fluxes and heating rates can be explained by

noting that, in spite of the fact that LWP and IWP compare favorably with observations, there is simply too much cloud cover in the model. The frequency of occurrence of cloudy pixels is nearly 10% and 20% greater in the model as compared with TRMM observations for the 46-day and 2-11 June time periods, respectively (Tables 2 and 3).

5 Sensitivity Experiments

Having determined that clouds and water vapor tend to be overpredicted in the control version of the GCE used in the NASA Goddard MMF, with concomitant effects on the cloud-affected radiation, sensitivity tests are now performed to assess whether the discrepancies between simulated and observed fields can be traced to the specifics of the model grid configuration, uncertainties in the cloud microphysical parameterization, or a combination of both. In the interest of brevity, results are presented only for the 2-11 June case; this allows identification of the sources of model error that are more directly associated with post-monsoon onset convection. Two sets of experiments are performed; one in which the model grid spacing and domain size are modified, and another in which assumptions in the GCE cloud microphysical parameterization that govern the ice particle size distribution are changed. The first set of comparisons is aimed at determining whether the differences between GCE simulated clouds, precipitation, and radiation are due to coarse grid spacing and/or a small horizontal domain. Specifically, if changes in domain size can be shown to improve simulation results, this will be an indication that connection of CRM domains across GCM grid cells may lead to an improvement in modeled clouds and precipitation. The second set of comparisons is designed to obtain an order of magnitude estimate of the sensitivity of CRM simulations in the MMF framework to changes in the assumptions contained in the GCE cloud microphysical parameterization. If the sensitivity to changes in cloud microphysical parameters is large, this result will provide motivation to further quantify and mitigate uncertainty in these parameters.

5.1 Grid Size and Spacing

In the cloud resolving modeling community, four kilometer grid spacing is generally accepted as an upper limit to what may legitimately be called “cloud resolving” (Weisman et al. 1997); at grid lengths greater than four kilometers, it is assumed that some manner of subgrid cumulus parameterization will be necessary to represent convection in the model.³ Since the CRMs used in current MMFs employ a four kilometer grid spacing, it is appropriate to consider whether the CRM results presented in this study are sensitive to changes in grid spacing. To this end, an experiment is performed in which the model grid spacing is halved to two kilometers (hereafter, 2KM) while the model domain is held fixed at 256 kilometers. While many studies have documented a sensitivity of simulations of cumulus cloud structure to changes in grid spacing, the role of mesoscale circulations in the evolution of deep convective systems, particularly those that form in the presence of vertical shear, has also been well documented (Fovell and Ogura 1989). Arguably, a 256 kilometer cyclic horizontal domain may be too small to allow development of realistic mesoscale circulations, hence a simulation is performed in which the model grid is extended to 1024 km in length (hereafter, 1024KM).

Results from 2KM and 1024KM simulations are presented along with TRMM observations of the 2-11 June time period in figure 7. Comparison of PWV between control, 2KM, and 1024KM simulations (Fig. 7a) reveals that a change in the grid spacing and grid length has a slight effect on the increase of moisture in the model. The fact that the 1024KM simulation more closely matches the TRMM observations than either the 2KM or control is likely due to a combination of two factors: (1) a larger domain provides a greater area over which compensating subsidence can act, and (2) in a larger domain, convective systems must

³Indeed, following recent case studies of turbulent structures in convection (Bryan et al. 2003) and vertically propagating gravity waves (Lane and Knievel 2005), it is now generally agreed that grid spacing on the order of 100 meters is necessary to appropriately resolve deep convection, with even smaller grid spacings required for the realistic simulation of trade-wind cumulus and stratocumulus clouds. Due to computational limitations, MMFs cannot be run with CRM grid spacings much less than 4 kilometers; the two kilometer grid spacing used here is assumed to be marginally computationally feasible.

travel a greater distance before encountering a cyclic boundary. Some sensitivity to both grid spacing and length is also observed in the LWP and IWP histograms (Figs. 7b and 7c), however, as indicated in the histograms and IAD values (Table 7), 2KM and 1024KM simulations provide little improvement over the control. Comparison of cloud fraction between TRMM, control, 2KM, and 1024KM (Table 3) indicates that the large-domain simulation produces a total cloud fraction that is more consistent with TRMM observations, while cloud cover in 2KM is even greater than in the control simulation. Neither simulation produces a more realistic precipitating fraction than the control; in fact, the large-domain simulation produces quite sparse precipitation compared with both the control and 2KM. The precipitation rate (Fig. 7d) itself is sensitive to changes in grid geometry, though neither the 2KM nor the 1024KM simulation provide a clear improvement over the control. The most notable difference is the reduction in large precipitation rates in 2KM, however, this improvement is offset by an anomalous increase in the moderate ($10\text{--}20 \text{ mm h}^{-1}$) rain rates. 1024KM exhibits precipitation rates that are even larger on average than the control simulation, and a precipitating fraction that is one third of what is observed by TRMM. Examination of the longwave radiative forcing (Fig. 7e) indicates that the 2KM simulation appears to provide a better fit to the observations than either the control or 1024KM, however, the model is still biased toward longwave forcing values that are too large compared with TRMM. Improvement in 2KM over the control may be due to a reduction in the occurrence of large LWP values and concurrent reduction of liquid cloud longwave absorption. Both simulations produce a shortwave cloud radiative forcing PDF (Fig. 7f) that is a better match to observations than the control; specifically, 1024KM exhibits less than 10% difference from the observed PDF (Table 7), with a marked reduction in the highest shortwave radiative forcing values. Comparison of the long and shortwave heating rates (Figs. 7g–7l) produced by 2KM and 1024KM with the control simulation reveal only slight differences in the simulated results, and it appears that changes to grid spacing and domain size do not affect the overall distribution of low, middle, and high clouds in the model.

Overall, neither an increase in grid resolution nor an increase in grid domain greatly improves the veracity of simulated results, though comparison of the total IAD values for each simulation (Table 7) indicates that both experiments provided a slight overall improvement. Comparison of rain rates and cloud fraction appears to indicate that 2KM may more effectively simulate stratiform rain than the control, but overproduces it—perhaps due to use of a small model domain. In contrast, the large domain of 1024KM produces improved cloud fraction, but, as in the control, generates too much convective precipitation relative to stratiform. Because recent studies have indicated that grid spacing of less than 500 meters may be required for proper simulation of deep convection (Bryan et al. 2003, W. Cotton, personal communication), an experiment was performed in which the grid spacing was reduced to 250 m (results not shown). This simulation produced results that were generally comparable to those of the 2KM run, with the exception of the shortwave cloud radiative forcing, which exhibited a significant bias toward large values.

5.2 Sensitivity Experiments: Cloud Microphysical Parameters

As was noted in section 3, the particle size distribution of precipitating ice species (snow and graupel) in the GCE cloud microphysical scheme is assumed to be exponential, with the number of particles of a given diameter D parameterized for condensate species x as

$$N_x(D) = N_{0x} \exp(-\lambda_x D), \quad (1)$$

where N_{0x} is the slope-intercept and the slope λ_x can be written

$$\lambda_x = \left(\frac{\pi \rho_x N_{0x}}{\rho_l q_x} \right)^{\frac{1}{4}}, \quad (2)$$

where ρ_x is the particle density. In addition, the fall velocity of precipitating particles is typically expressed as a power relationship (Mitchell 1996)

$$V_x(D) = a_x D^{b_x}. \quad (3)$$

Note that a_x , b_x , N_{0x} and ρ_x are constant parameters that must be specified in advance. These parameters, along with assumed-constant collision/collection efficiencies for the interactions between each pair of species, govern to some extent the properties of clouds, precipitation, and water vapor simulated by the model. To rigorously assess the sensitivity of the full set of cloud microphysical parameters used in the GCE, each parameter should be systematically varied while the others are held constant. Though such an exercise is beyond the scope of this study, it is useful to obtain an order of magnitude estimate of the effect of changes to these parameters on simulated output. Hence, in this section the uncertainty of simulated output to changes in cloud microphysical parameters is estimated by considering two scenarios; one in which precipitating ice particles are uniformly assumed to be small compared to the control, and one in which ice particles are assumed to be large. In the small-ice case (hereafter SMICE), the slope intercepts of the snow and graupel ice particle size distributions are increased, while the collection efficiencies and fall speed parameters are decreased. In the large-ice case (hereafter LGICE), the opposite is done; collection efficiencies and fall speed parameters are increased, while snow and graupel slope intercepts are decreased. In each experiment, perturbations are applied within the range of values reported in the literature (Locatelli and Hobbs 1974, Mitchell 1996), and results are compared with TRMM retrievals as in sections 4 and 5.1.

Histograms of fields output from both SMICE and LGICE simulations are presented in figure 8. The PDFs of precipitable water vapor from each case (Fig. 8a) indicate that, in contrast to the domain and grid spacing perturbation experiments, changes to ice microphysical parameters have a relatively large effect on the accumulated water vapor. The SMICE simulation exhibits an even larger moist bias than the control, with mean PWV of 75.56 kg m⁻² and IAD of 1.90, while the LGICE simulation produces an improved match to TRMM retrievals, with mean PWV of 70.82 kg m⁻² and IAD of 1.60 (Table 7). It is interesting that

both perturbation experiments yield better fits to the observed PDFs of LWP, IWP, and precipitation rate than the control, though close inspection of the IWP histograms reveals a tendency toward thicker ice clouds in the LGICE case as compared to SMICE (Fig. 8d). Examination of the LW cloud radiative forcing (Fig. 8e) reveals that the SMICE simulation yields nearly identical results as in the control, while the LGICE run produces a better fit at the larger values. This comes at the expense of a poorer fit to the observed PDF at small values; hence the IAD is nearly identical between SMICE and LGICE runs. Both SMICE and LGICE simulations produce a nearly identical tail in the shortwave radiative forcing (Fig. 8f), though the LGICE simulation produces a closer fit at smaller values, and thus the overall IAD is smaller (Table 7). It is interesting to note that both SMICE and LGICE produce a better fit to the observations than the control, primarily through a visible reduction in the number of SW radiative forcing values less than -400 W m^{-2} . This is reflected in the fact that the means of SMICE and LGICE simulations are closer to that of TRMM, and the IAD values are smaller than in the control. In contrast to the cloud, precipitation, vapor, and radiative forcing, the cloud-affected heating rates do not differ much between the control simulation, SMICE, and LGICE. Aside from a slightly better agreement in the mid-level shortwave radiative heating for LGICE, the histograms are nearly identical, as are the means and IAD values. As was the case for the control and grid sensitivity experiments, changes to cloud microphysical parameters appear not to affect the simulated distribution of low, middle, and high clouds, nor to mitigate the biases in cloud-affected radiative heating. The fact that the cloud fraction was similar between SMICE, LGICE, and control indicates that a combination of a larger grid with proper specification of cloud microphysical parameters may be needed for improved simulation of clouds and precipitation.

6 Discussion

6.1 Integrated Absolute Difference

Though most of the results presented in sections 4 and 5 involved visual inspection of PDF structures, the utility of the mean and integrated absolute difference as quantitative metrics in the statistical comparison was also clearly demonstrated. Use of these statistics becomes increasingly valuable as the work transitions from comparison to assimilation; in particular, given appropriate specification of error statistics, mean and IAD provide different pieces of information that together can be used to constrain uncertainty in the model. The primary drawback to the use of any integrated measure of difference between histograms lies in the discretization of the histogram into bins; if bin sizes are defined too small, spurious differences in histograms over small ranges will lead to consistently large IAD. By contrast, if bin sizes are set too large, the comparison may neglect important characteristics of either the observed or simulated PDF; e.g., secondary modes. The method that perhaps makes the most intuitive sense is to let the bin size depend on the error in the observations themselves. In this way, a measure of the observation uncertainty explicitly enters the process, and there is less chance that there will be noise in the comparison. In addition, it is unlikely that statistically significant structures can exist in the observed PDFs at scales that are smaller than, or comparable to the error magnitude.

6.2 CRM Feedback to the Large Scales

This study has considered the effect of changes in grid configuration and cloud microphysical parameters on the simulated histograms, however, an equally important consideration is the feedback of the CRM temperature and moisture tendencies to the large scale. This feedback can be computed following Khairoutdinov et al. (2005) as

$$\left[\frac{\partial \Phi}{\partial t} \right]_{CRM} = \frac{\overline{\Phi}^{n+1} - \Phi_{LS}}{\Delta t_{LS}}, \quad (4)$$

where Φ_{LS} is the variable of interest on the GCM grid and $\bar{\Phi}^{n+1}$ is the accumulated CRM tendency over the GCM timestep Δt_{LS} . In the current case, no feedback from the CRM to the large scales is allowed, and the CRM “forcing” is computed at half-hour intervals as

$$\left[\frac{\partial \Phi}{\partial t} \right]_{CRM} = \left[\frac{\partial \Phi}{\partial t} \right]_{total} - \left[\frac{\partial \Phi}{\partial t} \right]_{LS}, \quad (5)$$

where $\left[\frac{\partial \Phi}{\partial t} \right]_{total}$ is the total change in Φ over the half-hour interval, and $\left[\frac{\partial \Phi}{\partial t} \right]_{LS}$ is the tendency of Φ obtained from the observed large-scale forcing fields.

Total domain mean temperature and water vapor feedbacks are computed from equation (5) over the 2-11 June time period for the each simulation. The CRM temperature response (Fig. 9a) exhibits a structure that is nearly identical across all model configurations, with heating maximized in the middle troposphere and cooling observed near the surface and above 15 km. The vertical structure closely resembles superposition of a strong convective heating profile on a much weaker stratiform mode, and reinforces the notion that the model generates too much convective precipitation relative to stratiform. Cooling above 15 km, combined with heating between 12 and 15 km is consistent with the existence of thick cirrus anvils in all simulations. It is interesting to note that differences in the temperature feedback are greatest in the upper troposphere, a reflection of the vertically integrated effect of changes in cloud properties and distribution between different simulations. Though comparison with the temperature feedback reveals greater variability in the water vapor response, the CRM water vapor feedback (Fig. 9b) exhibits moistening in the lower and middle troposphere in all simulations. The moistening is greatest below 5 km, where evaporation of precipitation is occurring, and it can be seen that differences between simulations generally reflect the differences in the mean precipitable water vapor during the June time period (Table 6).

6.3 Role of Large Scale Forcing

The results presented in sections 4 and 5 above indicate that the GCE appears to produce an overactive response to large scale forcing, generating too much cloud and water vapor during periods of active convection. Though the large-scale forcing datasets used in this study have been extensively evaluated, and comparisons between PWV derived from the analysis and from TRMM demonstrate close agreement, it is fitting to examine how simulation results might change if the forcing were reduced. In this case, it is assumed that all sources of forcing are overestimated, and a constant correction factor is applied in the model equations so that the forcing magnitude is reduced to 80%. The control simulation is then re-run with the adjusted forcing, and fields are compared with the original control simulation. The results of reducing the large-scale forcing are in general comparable with the results from the unadjusted control simulation, with a few notable exceptions. In particular, the simulated precipitable water vapor, precipitation rate, LWP, and IWP (Fig. 10) are decreased relative to the control to values that are comparable with that observed in the LGICE simulation (Fig. 8). By contrast, the cloud-affected long and shortwave fluxes and heating rates are unaffected (results not shown). The implication is that, while a reduction in forcing appears to have led to a reduction in convective intensity, as measured by the precipitation rate, cloud thickness, and accumulation of water vapor, it does not appear that a bias in the large scale forcing fields is necessarily the reason for the discrepancies observed between simulation and observations.

7 Summary and Conclusions

In this study, clouds, precipitation, and long and shortwave radiation simulated by a cloud-resolving model were evaluated against retrievals from TRMM. The two-dimensional NASA Goddard cumulus ensemble cloud resolving model was run in an identical configuration to that used in the NASA Goddard multiscale modeling framework with a relatively small

Cotton 2004) has recently been included in the GCE model, and the radiative transfer of Gabriel et al. (2001) and Stephens et al. (2001) has also been implemented to provide more accurate treatment of the interaction of radiation with the RAMS cloud and ice condensate species. Tests of the RAMS microphysical scheme and Gabriel-Stephens radiation code are underway, and comparison of simulation results with TRMM observations is left for a future study. Finally, it is reasonable to expect that the role of cloud microphysical parameters might change when CRM simulations are run over different geographical regions and for different seasons. Further comparison of CRM results vs. observations will be necessary to determine whether this is, indeed, the case.

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Table 1: Observations used in the GCE comparison study, along with their units and error estimates.

Observation	Units	Error
Precipitable Water Vapor	$kg \cdot m^{-2}$	10 %
Precipitation Rate	$mm \cdot hr^{-1}$	25 %
Liquid Water Path	$kg \cdot m^{-2}$	40 %
Ice Water Path	$kg \cdot m^{-2}$	100 %
TOA LW Cloud Forcing	$W \cdot m^{-2}$	20 %
TOA SW Cloud Forcing	$W \cdot m^{-2}$	40 %
Three-Layer All-Sky LW and SW Radiative Cooling/Heating	$K \cdot hr^{-1}$	50 %
Three-Layer Clear-Sky LW and SW Radiative Cooling/Heating	$K \cdot hr^{-1}$	15 %

Table 2: Cloudy and precipitating fraction from TRMM and each of the GCE simulations over the entire 46-day integration period.

	TRMM	CTRL	SMICE	LGICE	$\Delta x = 2$ km	1000 km
Cloudy	0.42	0.51	0.54	0.49	0.53	0.43
Precipitating	0.13	0.05	0.05	0.05	0.05	0.04

Table 3: Cloudy and precipitating fraction from TRMM and each of the GCE simulations over the 2-11 June 1998 time period.

	TRMM	CTRL	SMICE	LGICE	$\Delta x = 2$ km	1000 km
Cloudy	0.60	0.80	0.85	0.79	0.82	0.63
Precipitating	0.16	0.09	0.10	0.10	0.10	0.06

Table 4: Mean values of each of the simulated and observed fields over the entire 46-day integration period.

Observation	TRMM	CTRL	SMICE	LGICE	$\Delta x = 2$ km	1000 km
PWV	55.21	61.94	63.05	60.58	62.07	60.80
LWP	0.09	0.57	0.51	0.53	0.49	0.63
IWP	0.14	0.56	0.40	0.52	0.43	0.50
Precip Rate	0.23	6.51	5.29	5.77	5.54	7.87
LW CRF	94.90	84.67	82.07	75.81	81.89	93.43
SW CRF	-90.66	-158.71	-145.43	-139.09	-142.25	-143.13
Low Qlw	0.48	1.10	1.12	1.03	1.00	0.99
Mid Qlw	-0.23	0.82	0.90	0.83	0.75	0.80
Upper Qlw	0.99	0.49	0.48	0.40	0.56	0.82
Low Qsw	-2.19	-2.91	-2.88	-2.92	-3.00	-2.86
Mid Qsw	-2.74	-3.67	-3.62	-3.61	-3.66	-3.49
Upper Qsw	-2.08	-3.74	-3.71	-3.53	-3.61	-3.74

Table 5: Integrated absolute difference between PDFs generated from each GCE simulation and those observed by TRMM over the full 46-day integration.

Observation	CTRL	SMICE	LGICE	$\Delta x = 2$ km	1024 km
PWV	1.28	1.27	1.18	1.33	1.15
LWP	0.07	0.08	0.08	0.09	0.17
IWP	0.09	0.07	0.07	0.07	0.07
Precip Rate	0.32	0.29	0.34	0.27	0.45
LW CRF	0.42	0.39	0.45	0.43	0.43
SW CRF	0.39	0.34	0.31	0.33	0.26
Low Qlw	1.35	1.34	1.24	1.35	1.27
Mid Qlw	1.06	1.12	0.99	1.03	1.07
Upper Qlw	0.77	0.85	0.81	0.73	0.69
Low Qsw	0.96	0.98	0.97	0.99	0.89
Mid Qsw	1.09	1.11	1.06	1.08	0.97
Upper Qsw	1.19	1.39	1.08	1.27	1.37
Total	9.00	9.25	8.60	8.96	8.79

Table 6: Mean values of each of the simulated and observed fields over the 2-11 June 1998 time period.

Observation	TRMM	CTRL	SMICE	LGICE	$\Delta x = 2$ km	1024 km
PWV	58.83	73.63	75.56	70.82	73.61	72.51
LWP	0.17	0.72	0.71	0.62	0.72	0.73
IWP	0.28	0.84	0.47	0.66	0.66	0.73
Precip Rate	0.48	9.58	6.53	6.93	8.16	10.01
LW CRF	108.74	105.26	101.45	92.23	104.59	100.94
SW CRF	-26.53	-76.71	-65.76	-64.12	-70.83	-54.74
Low Qlw	0.42	1.24	1.29	1.08	1.02	0.96
Mid Qlw	-0.04	1.44	1.42	1.32	1.38	1.12
Upper Qlw	1.17	0.57	0.62	0.61	0.75	0.86
Low Qsw	-2.02	-2.71	-2.66	-2.80	-2.73	-2.70
Mid Qsw	-2.43	-3.00	-2.98	-3.06	-3.00	-2.95
Upper Qsw	-1.99	-2.94	-2.91	-2.80	-2.94	-2.87

Table 7: Integrated absolute difference between PDFs generated from each GCE simulation and those observed by TRMM over the 2-11 June 1998 time period.

Observation	CTRL	SMICE	LGICE	$\Delta x = 2$ km	1024 km
PWV	1.83	1.90	1.60	1.90	1.73
LWP	0.18	0.11	0.13	0.19	0.21
IWP	0.18	0.14	0.14	0.21	0.15
Precip Rate	0.55	0.37	0.40	0.49	0.53
LW CRF	0.45	0.48	0.49	0.44	0.68
SW CRF	0.30	0.28	0.24	0.30	0.17
Low Qlw	1.51	1.43	1.30	1.41	1.28
Mid Qlw	1.40	1.40	1.19	1.28	1.17
Upper Qlw	0.79	0.85	0.85	0.67	0.79
Low Qsw	1.56	1.53	1.55	1.55	1.41
Mid Qsw	1.35	1.34	1.31	1.33	1.14
Upper Qsw	1.20	1.56	0.83	1.46	1.43
Total	11.30	11.38	10.00	11.22	10.70

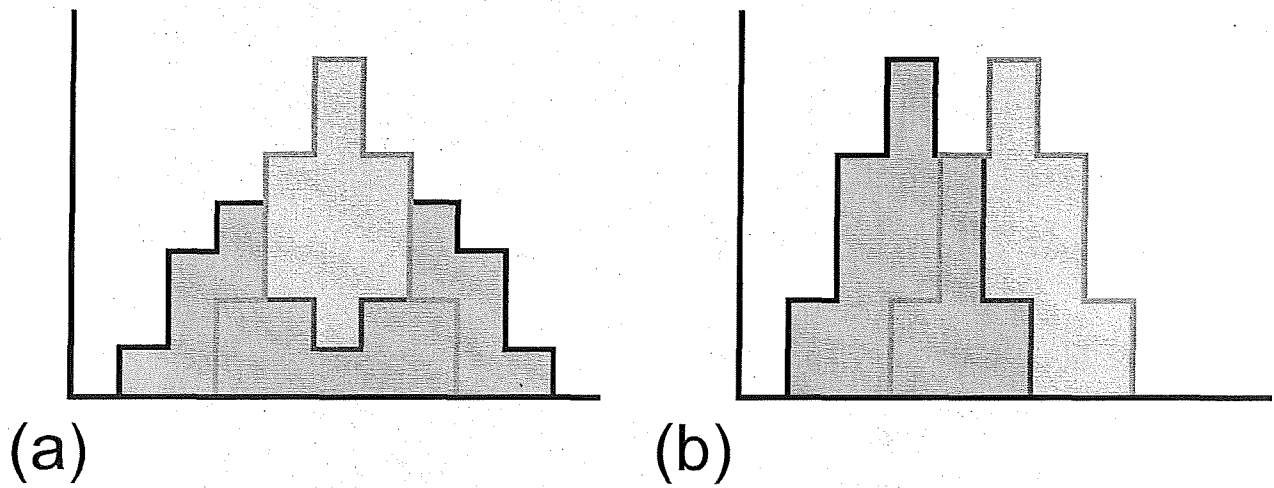


Figure 1: Illustration of the utility of both mean and integrated absolute difference in comparing PDFs between observation and model. PDFs for simulated observations are depicted in gray, while model results are depicted in black. Both figures have identical integrated absolute difference, but different means.

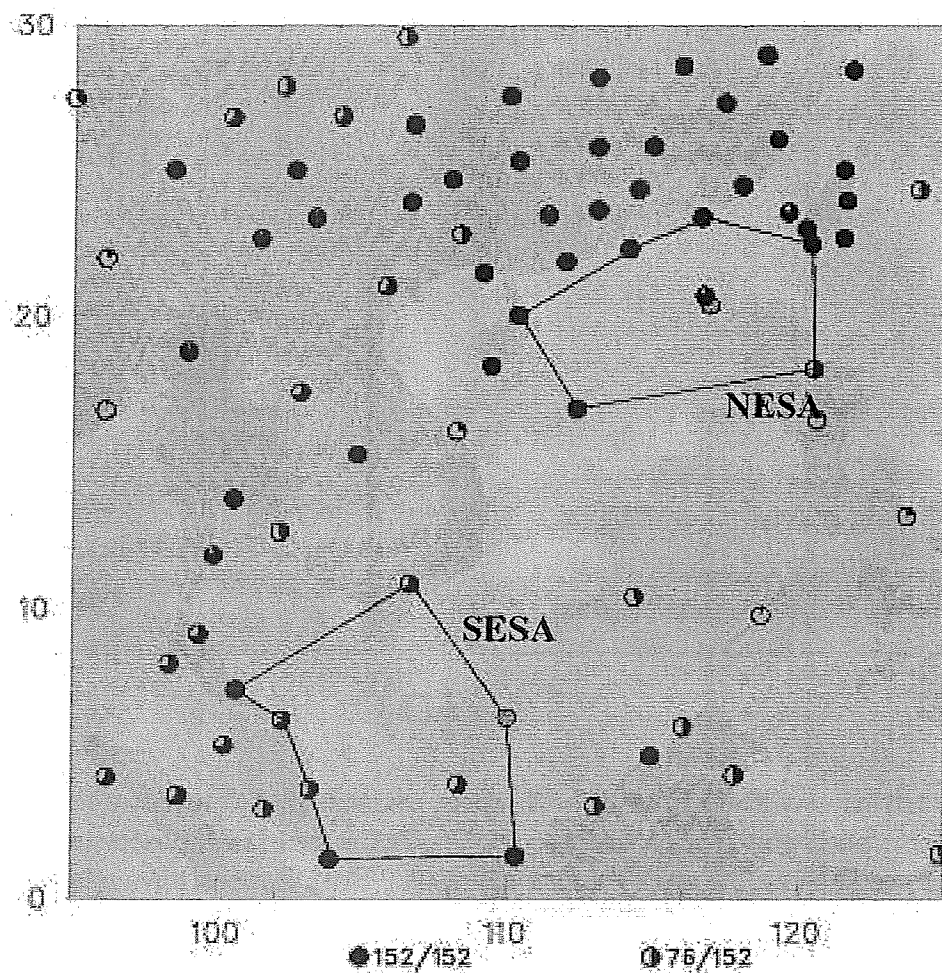


Figure 2: Depiction of the SCSMEX NESAs and SESAs with sounding sites marked in black and white circles; the portion of each circle that is filled corresponds to the fraction of the 46-day field campaign during which soundings were available four times daily. The locations of the dual-doppler network and *Kexue* #1 research vessel are depicted in the gray filled circles in the NESAs and SESAs regions, respectively. Figure obtained from http://tornado.atmos.colostate.edu/scsmex/nesa-averaged_v3b.html.

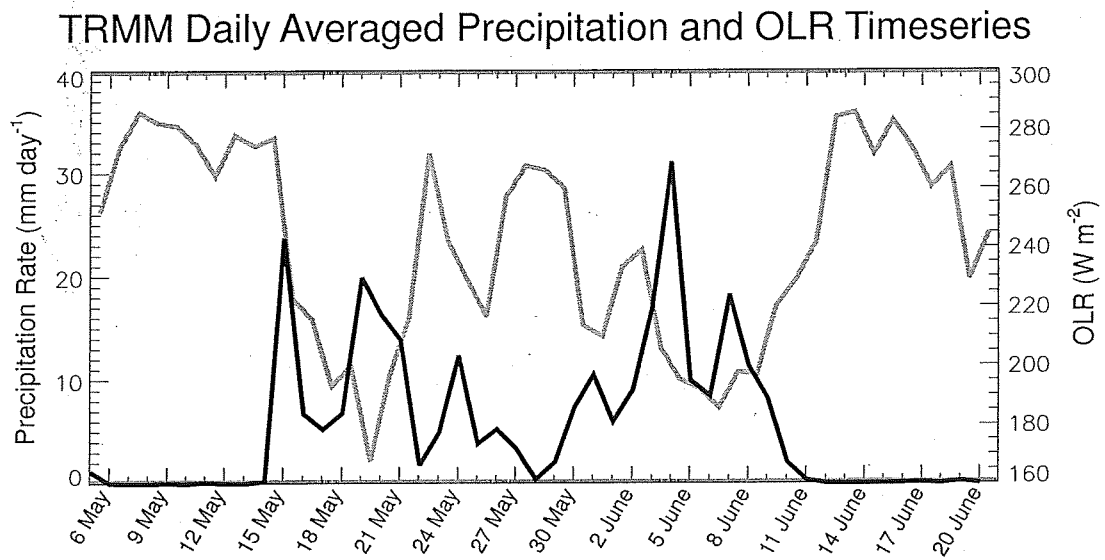


Figure 3: Depiction of the daily domain-averaged precipitation rate (black) and outgoing longwave radiative flux (gray) retrieved from TRMM over the 46-day SCSMEX intensive observation period.

Control, All Times, Matched to TRMM

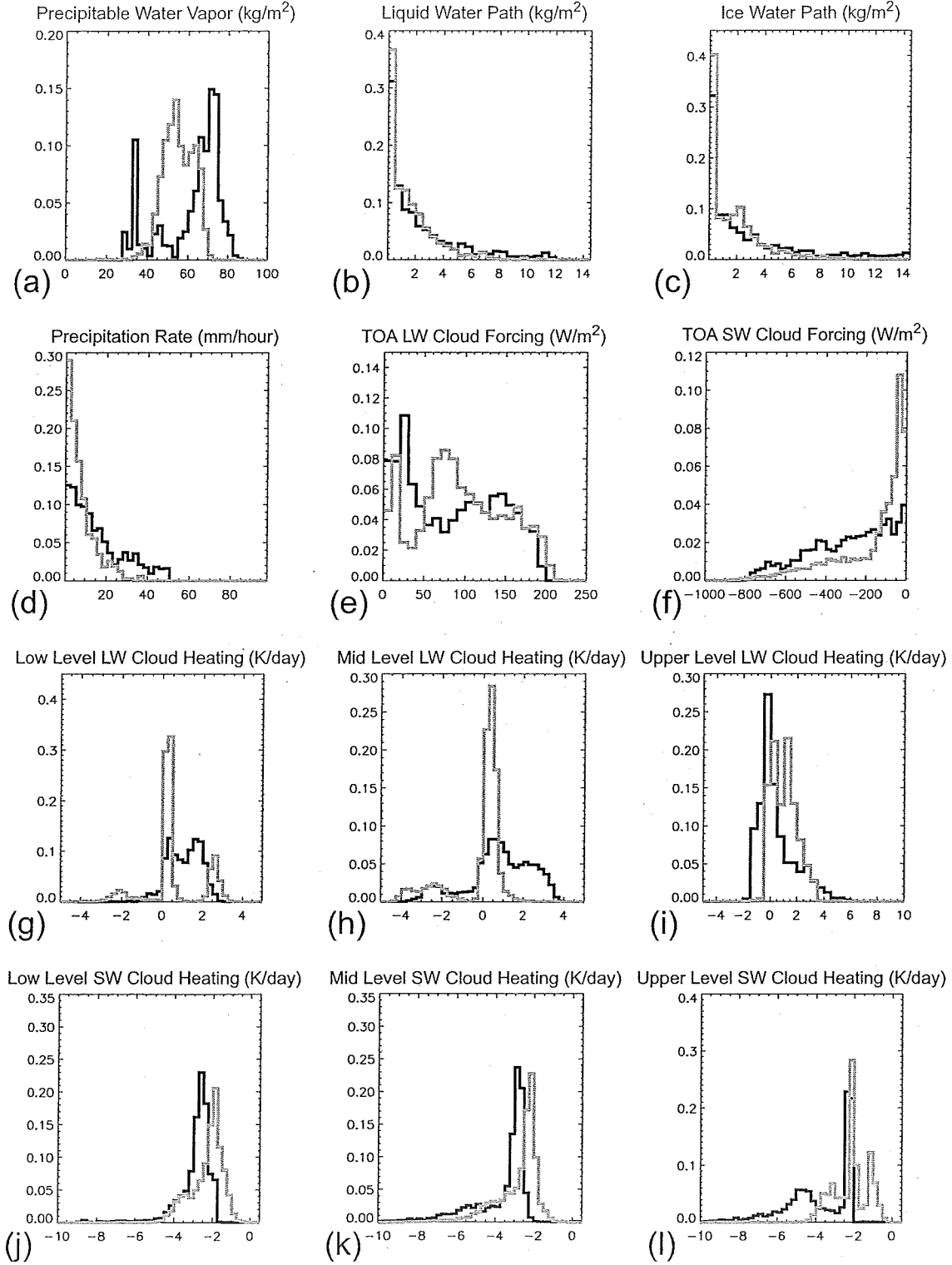


Figure 4: Histograms of TRMM observations (gray) and GCE simulated output (black) for the control simulation. Results are presented for the entire 46-day time period.

Control, 2-11 June, Matched to TRMM

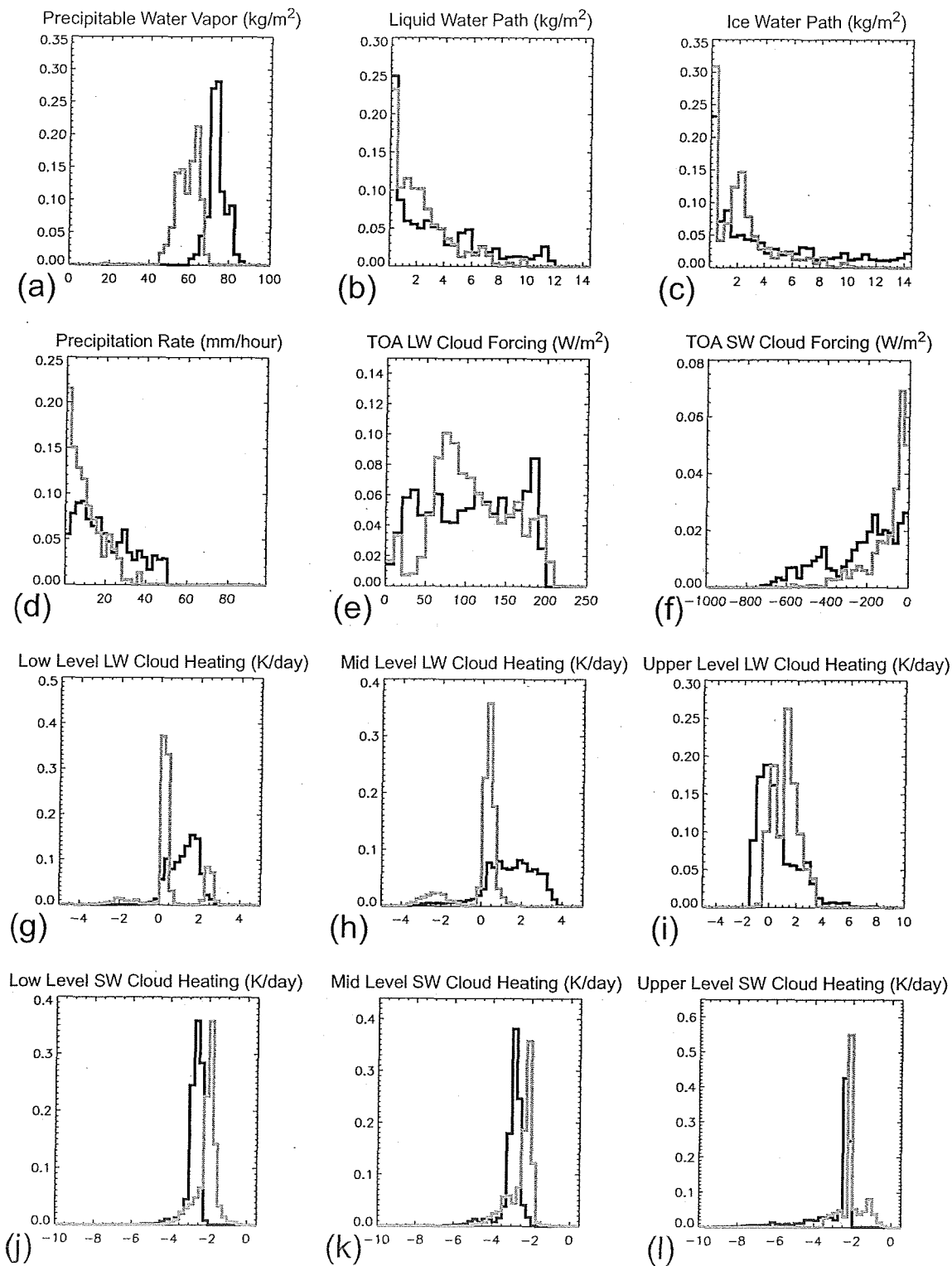


Figure 5: As in figure (4), but for the 2-11 June 1998 time period.

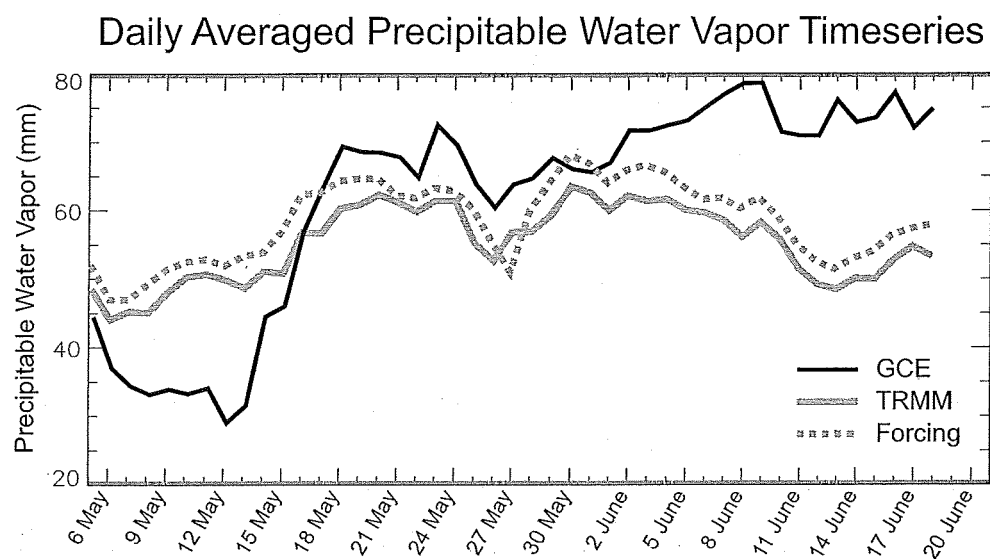


Figure 6: Timeseries of daily-mean precipitable water vapor from TRMM retrievals (gray solid line), SCSMEX forcing dataset (gray dashed line), and the GCE control simulation (black solid line).

Grid Comparison, 2-11 June, Matched to TRMM

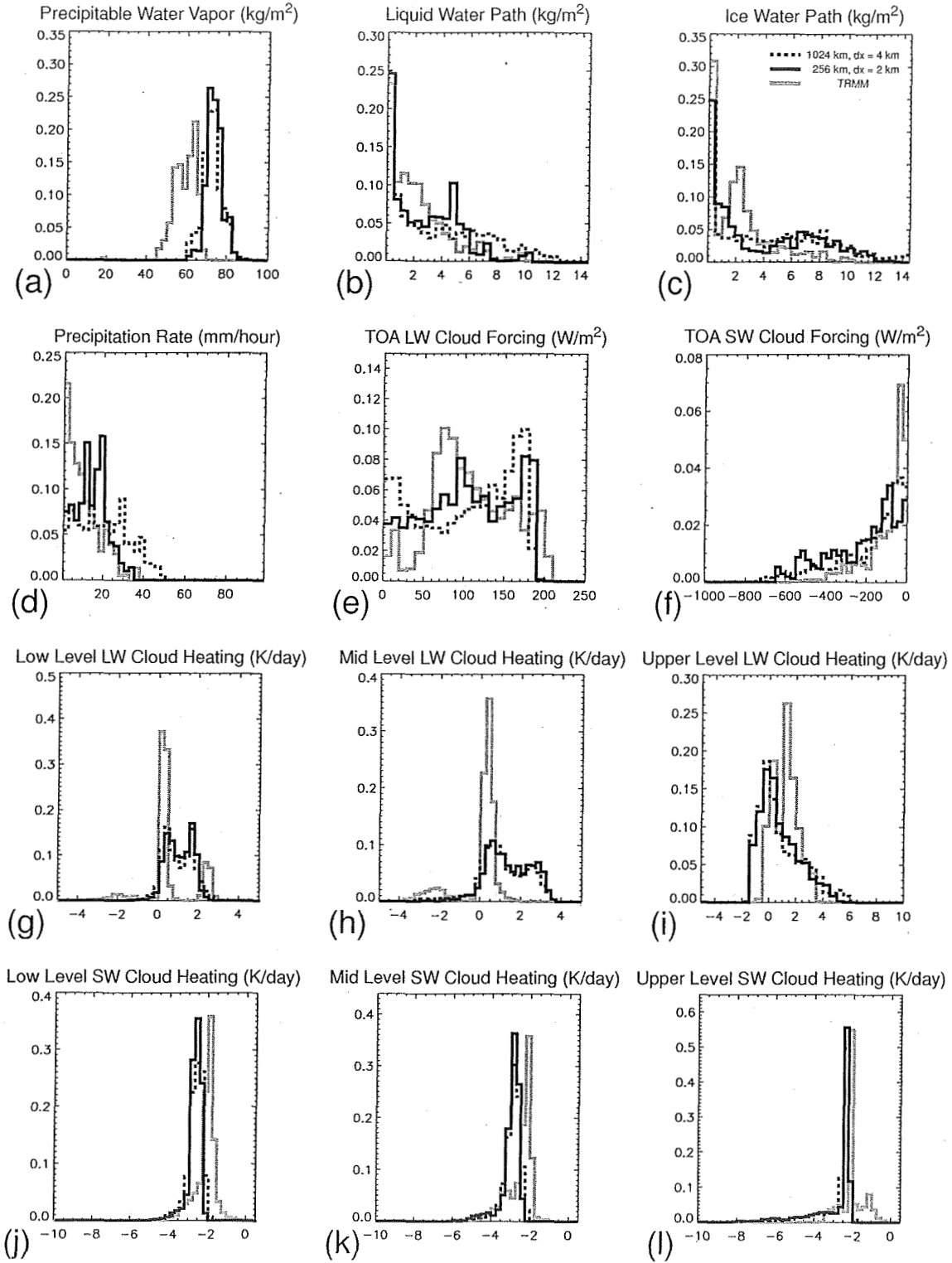


Figure 7: Comparison of TRMM observations (gray) with GCE simulations for runs with 256 km domain and 2 km grid spacing (black solid) and with 1024 km domain and 4 km grid spacing (black dashed). Results are presented for the 2-11 June 1998 time period.

Microphysics Comparison, 2-11 June, Matched to TRMM

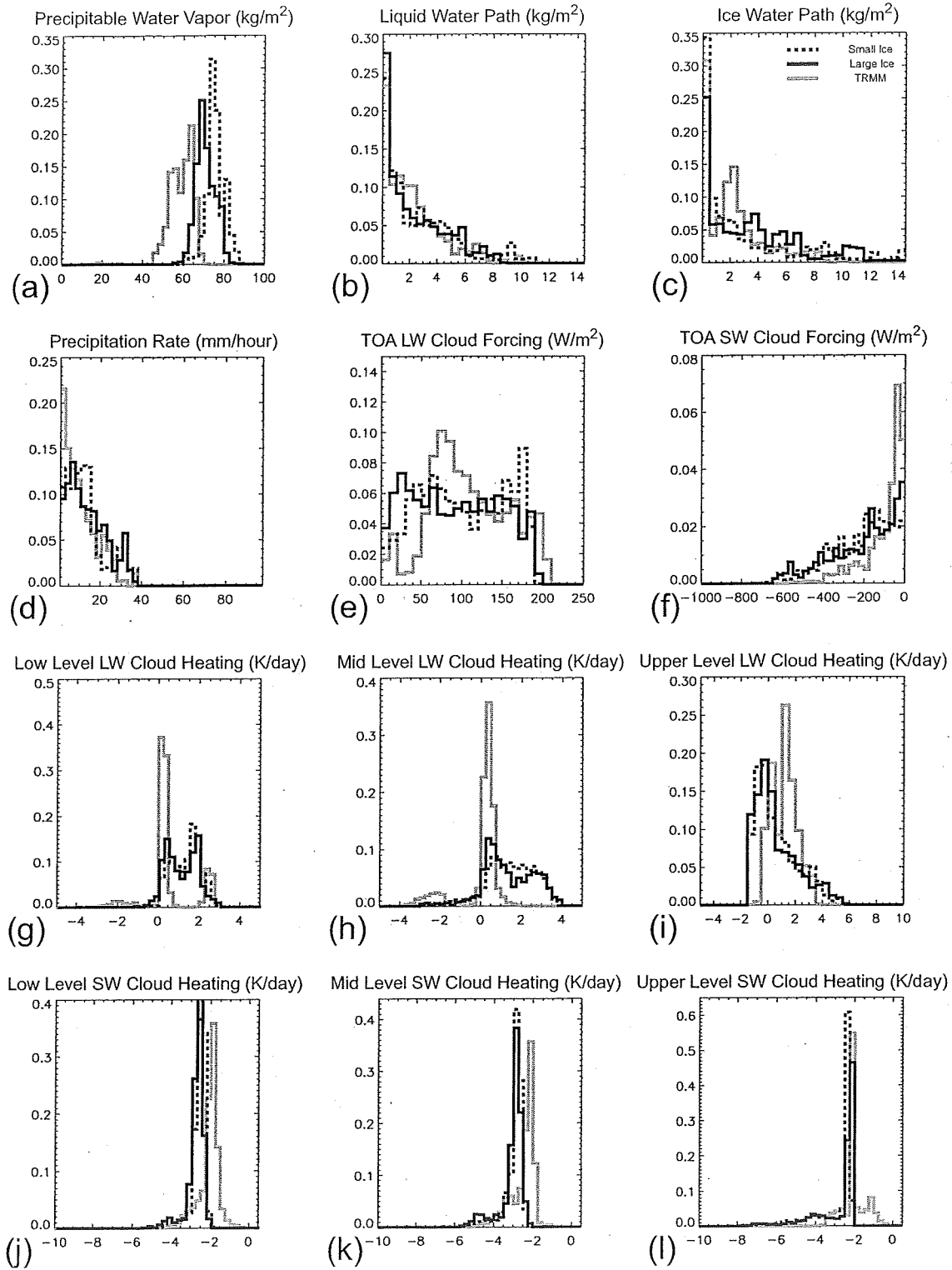
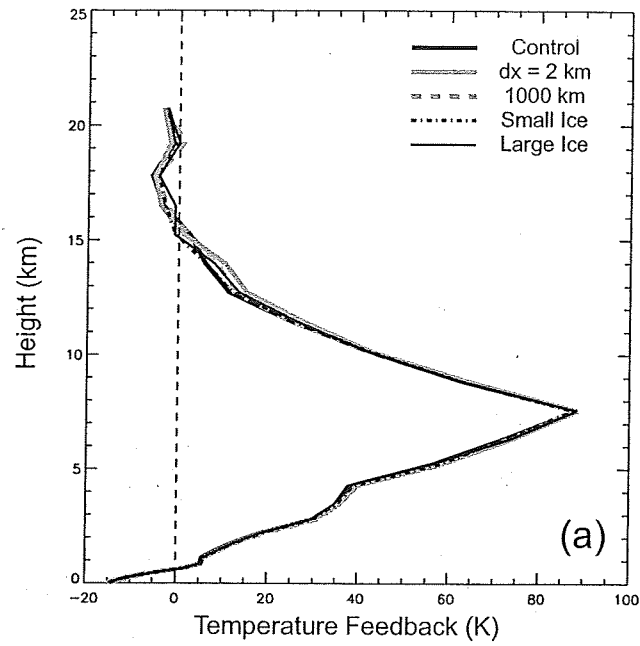


Figure 8: Comparison of TRMM observations (gray) with GCE simulations for runs with assumed small ice particles (black dashed) and with assumed large ice particles (black solid). Each simulation is run with 4 km grid spacing on an 256 km domain, and results are presented for the 2-11 June 1998 time period.

CRM Total Temperature Feedback, 2-11 June 1998



CRM Total Water Vapor Feedback, 2-11 June 1998

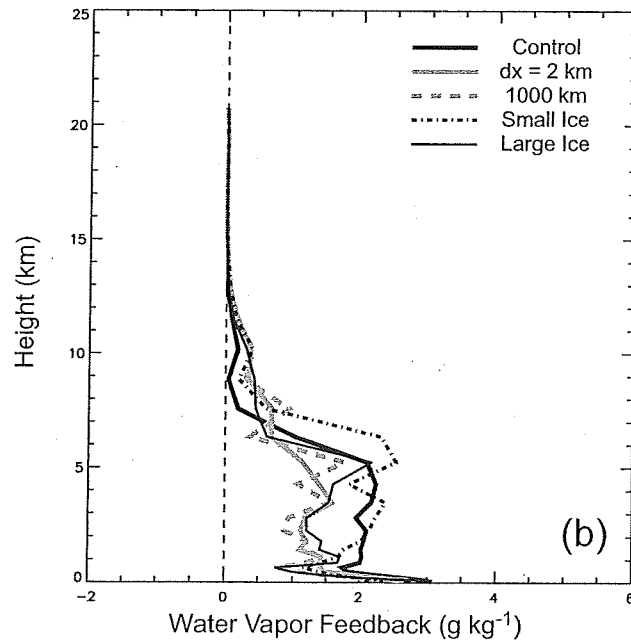


Figure 9: Accumulated GCE temperature (a) and water vapor (b) tendencies over the 2-11 June time period as a function of height, averaged in the horizontal.

Reduced Forcing, 2-11 June, Matched to TRMM

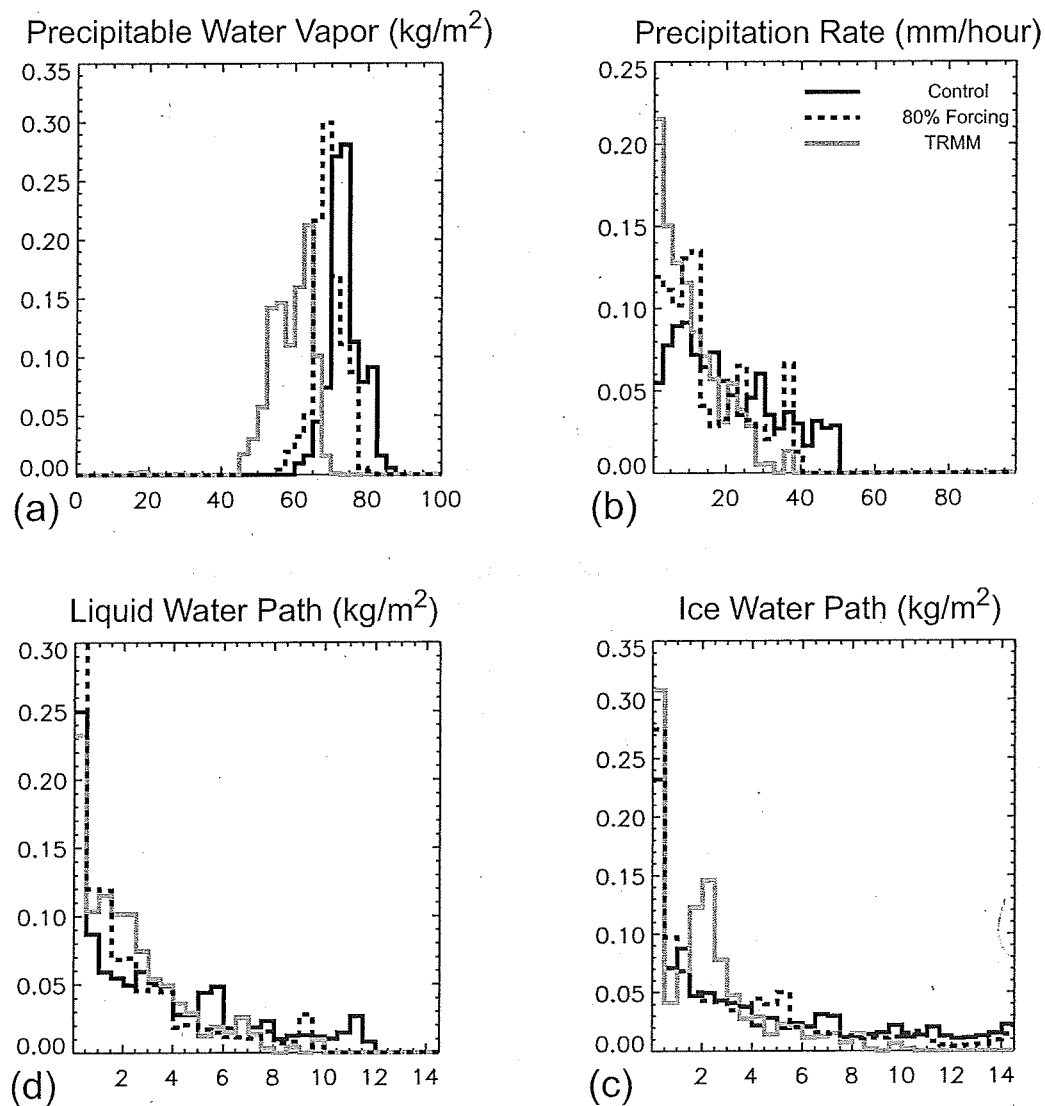


Figure 10: Comparison of TRMM retrievals (gray) with control simulation (black solid) and the control simulation forced with 80% large scale forcing fields (black dashed).